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准噶尔盆地盆1井西凹陷及周缘深层二叠系 超压形成机制及演化特征

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摘要: 为明确准噶尔盆地盆1井西凹陷及周缘深层二叠系超压形成机制及演化特征, 基于钻井、测井、实测地层压力等资料, 运用测井曲线组合分析法和交会图版法以及盆地模拟技术, 对超压成因机制及演化特征进行分析, 并定量表征目的层不同成因类型超压的贡献率。结果表明: ①研究区深层二叠系不同岩性地层超压成因具有明显差异, 其中风城组和下乌尔禾组烃源岩层超压成因主要为生烃和欠压实作用, 其中生烃作用占主导; 泥岩盖层超压成因主要为欠压实作用; 而风城组和下乌尔禾组储层超压成因主要为超压传递和欠压实作用。②研究区风城组烃源岩生烃增压自早二叠世开始, 现今达到最大, 不同构造部位增压大小差别较小, 烃源岩层超压主要分布于 39.43 ~ 49.16 MPa, 但其贡献率具有明显差异, 凸起区较大, 而凹陷区较小; 凸起区和凹陷区生烃增压对总超压贡献率分别为 84.49% ~ 94.41% 和 65% ~ 67.3%, 研究区下乌尔禾组烃源岩生烃增压大小与风城组具有相似的规律; 欠压实作用对泥岩盖层超压的贡献率一般为 100%。③研究区二叠系储层超压传递增压主要形成于晚侏罗世至早白垩世以及古近纪至今, 其贡献率在不同凸起区存在明显差异, 达巴松凸起和石西凸起分别为 21.86% ~ 23.35% 和 100%。厘清研究区深层、超深层超压的分布规律, 可以为盆地新区的进一步开发提供依据。

关键词: 深层; 超压成因; 演化特征; 定量评价; 盆1井西凹陷; 准噶尔盆地

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Formation mechanism and evolution characteristics of deep Permian overpressure in western Well Pen-1 Sag and its periphery, Junggar Basin

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Abstract: To clarify the formation mechanism and evolution characteristics of the deep Permian overpressure in western Well Pen-1 Sag and its periphery in Junggar Basin, based on the data of drilling, logging, and measured formation pressure, the formation

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mechanism and evolution characteristics of overpressure were comprehensively studied, and the contribution rate of different types of overpressure genesis in the target layer was quantitatively characterized by using the methods of analysis of well logging curve combination, cross plots, and basin simulation. The results show that ① the overpressure genesis of different lithologies of the deep Permian in the study area have obvious differences, among which the overpressure genesis of source rock of Fengcheng Formation and the Lower Wuerhe Formation are mainly hydrocarbon generation and disequilibrium compaction, with hydrocarbon generation dominating; the overpressure genesis of mudstone caprock is mainly disequilibrium compaction, while the overpressure genesis of reservoirs of the Fengcheng Formation and Lower Wuerhe Formation are mainly overpressure transfer and disequilibrium compaction. ② The overpressure induced by hydrocarbon generation of Fengcheng Formation in the study area has appeared since the Early Permian and reached the maximum at present, and there are little differences in the size of the pressurization among different tectonic positions; the overpressure of the source rock is mainly distributed in the range of 39.43-49.16 MPa, but its contribution rate has obvious differences, with larger salients and smaller sags, and the contribution rates of pressurization induced by hydrocarbon generation in salients and sags to the total overpressure are 84.49%-94.41% and 65%-67.3%, respectively. The pressurization induced by hydrocarbon generation in the source rock of the Lower Wuerhe Formation has a similar size to that of Fengcheng Formation. The contribution rate of disequilibrium compaction to the overpressure in the mudstone caprock is basically 100%. ③ Overpressure transfer pressurizations in the Permian reservoir of the study area are mainly formed from the Late Jurassic to Early Cretaceous and since the Paleoproterozoic to now, and their contribution rates vary significantly in different salient areas, with 21.86%-23.35% and 100% in Dabasong Salient and Shixi Salient, respectively. Clarifying the overpressure distribution law of deep and ultra-deep formations in the study area can provide a basis for the further development of the new basin.

Key words: deep formation; overpressure genesis; evolution characteristics; quantitative evaluation; western Well Pen-1 sag; Junggar Basin

随着油气勘探的深入,深层、超深层领域不断获得勘探突破^[1-4],例如塔里木盆地的博孜9井^[5]和准噶尔盆地玛湖28、石西16、石西18、高探1、夏盐26等重点探井的勘探发现^[6-7],预示着深层、超深层已成为中国油气勘探的重要领域^[8-9]。深层、超深层广泛发育超压,与中、浅层相比,超压发育强度更高^[6,10],并且由于深层一般都经历了多期构造运动^[11-12],超压对于深层油气的生成、运移和聚集的影响更为复杂^[13-14],因此研究超压的形成机制及演化特征对油气勘探开发具有重要意义。

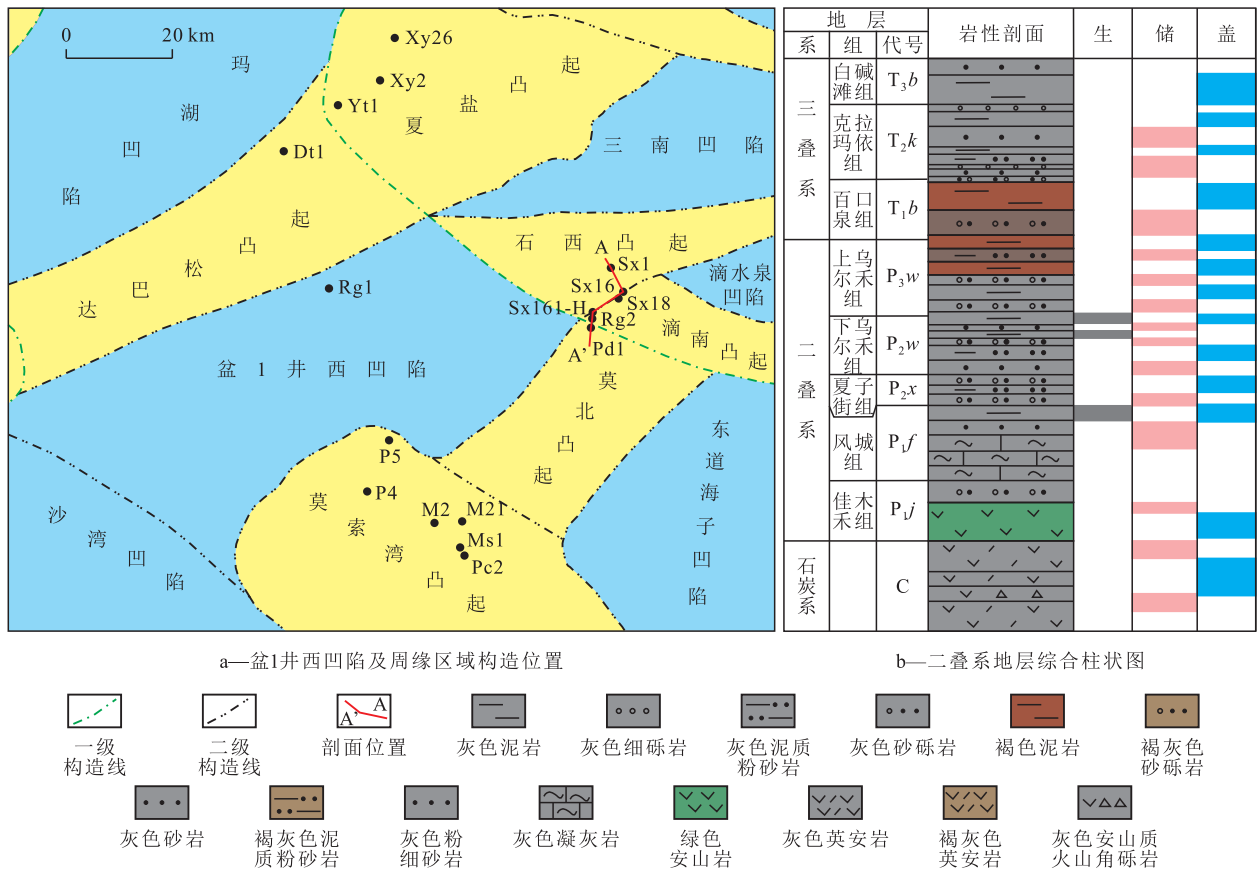
近些年来,不同学者对准噶尔盆地盆1井西凹陷及周缘侏罗系及以下层系的超压形成机制、演化特征及超压贡献率进行了研究^[14-20]。ZHANG等通过对准噶尔盆地中央凹陷中生界超压成因研究,认为其砂岩储层超压成因主要是超压传递作用,而泥岩超压成因主要与欠压实作用和化学压实作用有关,生烃作用有限^[21]。WANG等通过对盆地腹部侏罗系及以下层系的研究,认为烃源岩生烃和构造挤压作用是二叠系超压的主要成因,二叠系烃源岩生烃通过垂向压力传递导致侏罗系砂岩储层超压,欠压实作用不是该储层超压的主要原因^[22]。以往研究大多局限于侏罗系,对于研究区深层二叠系不同岩性地层超压形成机制研究较少,且对于二叠系超压的演化过程及其贡献率未见研究^[14-22]。笔者以准噶尔盆地盆1井西凹陷及周缘为研究对象,基于钻

井、测井、试油等资料,运用测井曲线组合分析法、交会图版法,结合实际地质分析,对研究区二叠系不同岩性地层超压成因进行研究,运用盆地数值模拟方法,对二叠系源、储、盖层在不同地质历史时期的地层压力演化进行定量恢复,厘清研究区不同构造部位超压的演化特征及各超压成因的贡献率,以期研究区油气勘探工作提供指导,也为同类型地质背景超压成因研究提供参考。

1 区域地质概况

准噶尔盆地位于中国新疆维吾尔自治区北部,是中国除塔里木盆地以外的第二大内陆盆地,为典型的古生代和中新生代多旋回叠合沉积盆地^[6,23]。根据盆地基底形态和后期构造演化作用,将准噶尔盆地分为乌伦古坳陷、陆梁隆起、西部隆起、中央坳陷、东部隆起和南缘冲断带6个一级构造单元^[24-25]。研究区位于中央坳陷的中北部,包括盆1井西凹陷、达巴松凸起、石西凸起、夏盐凸起、莫索湾凸起、莫北凸起及其环带,其西南部紧邻沙湾凹陷,西北部为玛湖凹陷,东部为东道海子凹陷^[26](图1a)。由于研究区地质资料受限,本次以达巴松凸起和石西凸起作为主要研究区域。

盆1井西凹陷及其周缘凸起主要是由海西期构造运动形成,石炭纪为快速沉降阶段,经历了2个断



a—盆1井西凹陷及周缘区域构造位置

b—二叠系地层综合柱状图

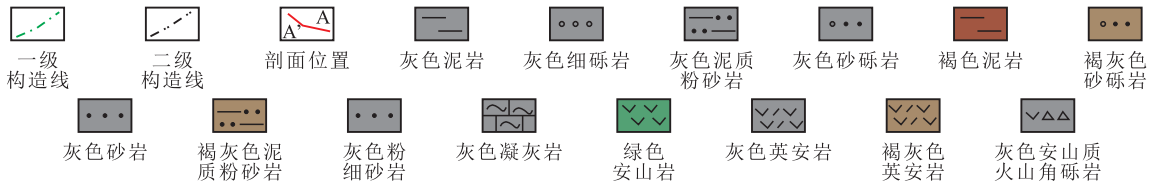


图1 准噶尔盆地盆1井西凹陷及周缘区域构造位置和二叠系地层综合柱状图

Fig.1 Comprehensive stratigraphic column of Permian formation and structural location of western Well Pen-1 Sag and its periphery, Junggar Basin

陷旋回^[22]。二叠世经历了佳木河组、风城组、夏子街组及上、下乌尔禾组的发育期,早二叠世为盆地的伸展阶段,中二叠世—三叠纪表现为沉降期,为前陆盆地发育阶段^[27],构造相对稳定^[28],其中中二叠世为快速沉降期,从晚二叠世—中三叠世沉降速度减慢,晚三叠世沉降缓慢^[27]。早—中侏罗世表现为快速沉降期,晚侏罗世构造抬升。自白垩纪以来,盆地一直沉降至今,为统一的拗陷阶段,构造活动强烈区由盆1井西凹陷及其周缘迁移至北天山山前拗陷,使早期形成的圈闭得以有效保存^[28]。研究区地层自下而上为石炭系、二叠系、三叠系、侏罗系、白垩系,发育石炭系、二叠系、三叠系、侏罗系等多套烃源岩层系,其中二叠系发育2套优质烃源岩层,分别为风城组(P₁f)和下乌尔禾组(P₂w),对盆1井西凹陷二叠系油气聚集贡献最大^[29-33](图1b)。风城组烃源岩干酪根类型主要为I—II₁型,从该烃源岩实测的总有机碳含量、生烃潜量、氢指数数据来看,其具有较高的有机质丰度特征;下乌尔禾组烃源岩干酪根类型主要为II₂—III型,有机质丰度较低,总体均为较好级别烃源岩^[24, 28, 34-36]。

研究区二叠系在纵向上发育佳木河组砾岩、风

城组砾岩、夏子街组和下乌尔禾组砂砾岩以及上乌尔禾组中下部厚层砾岩储层,发育风城组云质泥岩区域盖层、夏子街组上部盖层、下乌尔禾组泥岩区域盖层以及上乌尔禾组顶部的滨湖相泥岩局部盖层^[29, 31-32, 37]。其中,下乌尔禾组、夏子街组砂砾岩储层与下乌尔禾组泥岩区域性盖层以及风城组、佳木河组砾岩储层和风城组云质泥岩区域性盖层组成了良好的储-盖组合^[37-38],以风城组和下乌尔禾组储层为本次研究的重点层位。

2 现今实测地层压力分布特征

钻杆测试(DST)可以准确反映地层超压信息,由于研究区钻遇二叠系的井较少,本次共收集3口井9个深度点的实测地层压力数据。根据李伟等的超压划分方案^[39],将地层压力系数小于0.9定义为负压,地层压力系数为0.9~1.2定义为常压,地层压力系数大于1.2定义为超压,进而分别以地层压力系数为1.2,1.6和2.0作为界限,划分为弱超压、强超压和极强超压3个区域。研究区二叠系普遍存在超压,主要分布于17.61~47.41 MPa,不同构造部位超

压大小存在差异;其中,石西凸起地层超压分布于17.61~23.58 MPa,地层压力系数最大为1.53,位于弱超压区间;达巴松凸起地层超压分布于41.20~47.41 MPa,地层压力系数为1.8~2.0,最大为1.89,属于强超压区域范畴(图2),且达巴松凸起较石西凸起地层压力系数更高。

3 盆1井西凹陷超压形成机制

含油气盆地超压形成机制复杂,将其归类后主要有以下5种类型:欠压实作用、流体膨胀、成岩作用、构造挤压以及超压传递^[21,40-42]。本次运用测井曲线组合分析法和交会图版法进行超压成因

研究^[43]。

3.1 测井曲线组合分析法

沉积盆地超压判识最基本且较为可靠的方法是测井曲线组合分析法,这种方法能降低超压成因判断的不准确性^[44]。通常利用声波时差、密度、电阻率、中子孔隙度4条测井曲线,对超压成因进行综合判识。目前欠压实作用和生烃增压作用已成为含油气盆地两大主要的超压成因机制^[44-45]。

为了深入分析准噶尔盆地盆1井西凹陷及周缘的超压形成机制,统计研究区Dt1井、Sx18井泥岩的声波时差、密度、电阻率及中子孔隙度测井数据,通过密度测井资料对密度孔隙度进行计算,再结合密度和声波资料计算出声波孔隙度^[41]。

Dt1井侏罗系八道湾组底部(深度为3 882 m)泥岩测井曲线(图3)显示,泥岩声波时差明显高于上覆的三工河组,且其电阻率和密度有明显减小,中子孔隙度增大,表明出现异常高压。三叠系白碱滩组泥岩、克拉玛依组泥岩及二叠系西山窑组泥岩的中子孔隙度分别约为对应正常压实段的1~2倍,二叠系风城组储层测井解释孔隙度为6%~9%,下乌尔禾组储层测井解释孔隙度高达14.38%,均大于同等深度下正常压实孔隙度,且以原生孔隙为主,原生孔隙占总孔隙的比例为65%~100%^[6]。因此,Dt1井在深度大于3 882 m时表现出孔隙度异常高的欠压实特征。另外,下乌尔禾组上覆厚层灰色、深灰色泥岩以及砂质泥岩盖层,发育泥质细砂岩储层,沉积速率介于70~110 m/Ma;其快速沉积及岩性特征是形成欠压实作用的有利地质条件。综上所述,推断欠压实作用为达巴松凸起侏罗系八道湾组及下伏烃源岩层和储层超压的主要形成机制之一。

通过分析位于滴南凸起Sx18井的声波、密度、电阻率及中子孔隙度随深度的变化(图4)发现,当深度小于3 898 m时,泥岩以正常压实为主;当深度大于3 898 m时,声波时差和中子孔隙度出现正异常,电阻率出现负异常,密度无明显变化,四者反转不同步,表明在深度大于3 898 m时出现超压。八道湾组储层测井解释孔隙度为9.35%,二叠系泥岩、砂岩孔隙度及渗透率无实测,参照邻区玛湖凹陷风城组砂岩的孔隙度平均为5.44%、渗透率平均为0.1 mD,整体上属于特低孔特低渗透储层^[46],地层沉积速率介于3~7 m/Ma,沉积速率较低,不符合与高沉积速率和孔隙度异常高成因密切相关的欠压实超压。

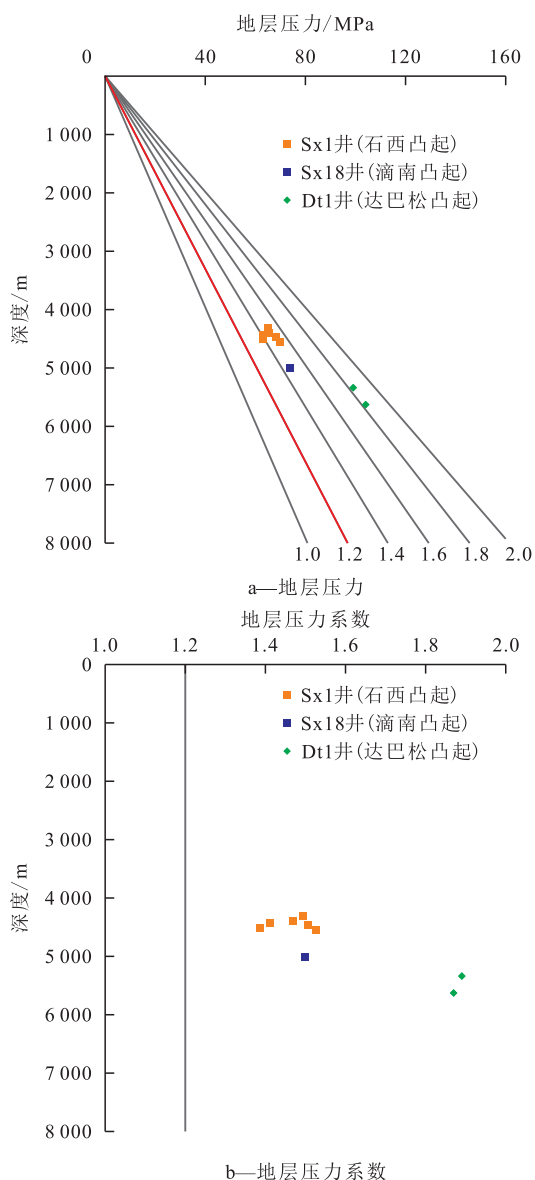


图2 准噶尔盆地盆1井西凹陷及周缘二叠系实测地层压力分布

Fig.2 Measured formation pressure distribution of Permian in western Well Pen-1 Sag and its periphery, Junggar Basin

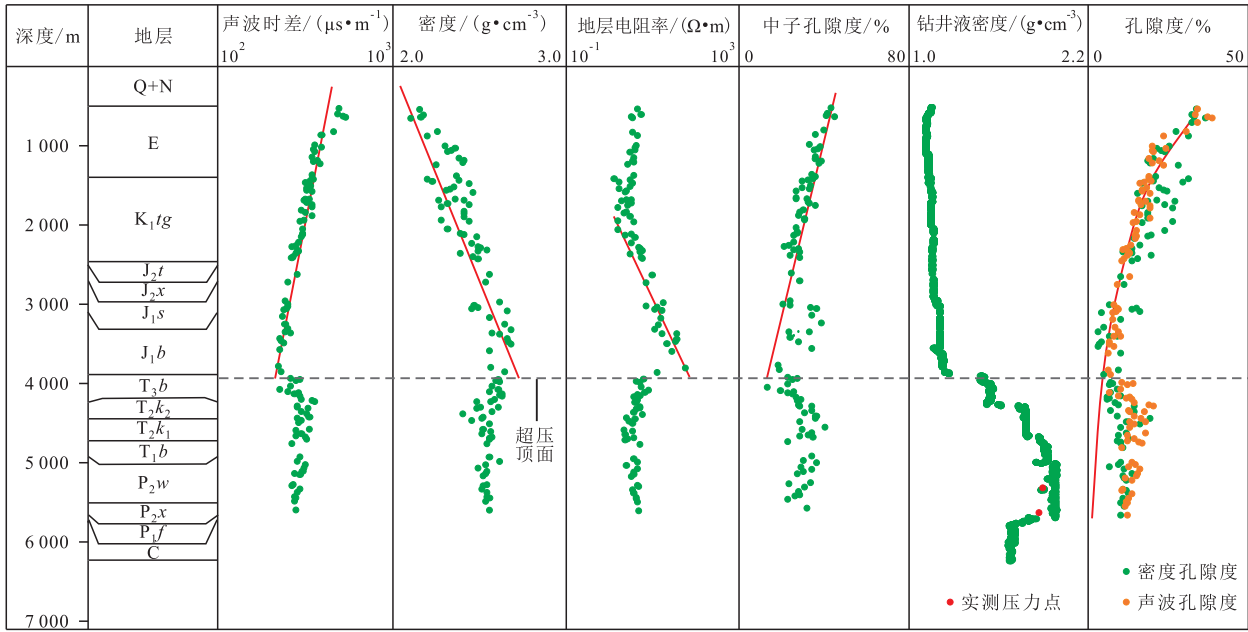


图3 Dt1井测井曲线组合分析
Fig.3 Analysis of well logging curve combination of Well Dt1

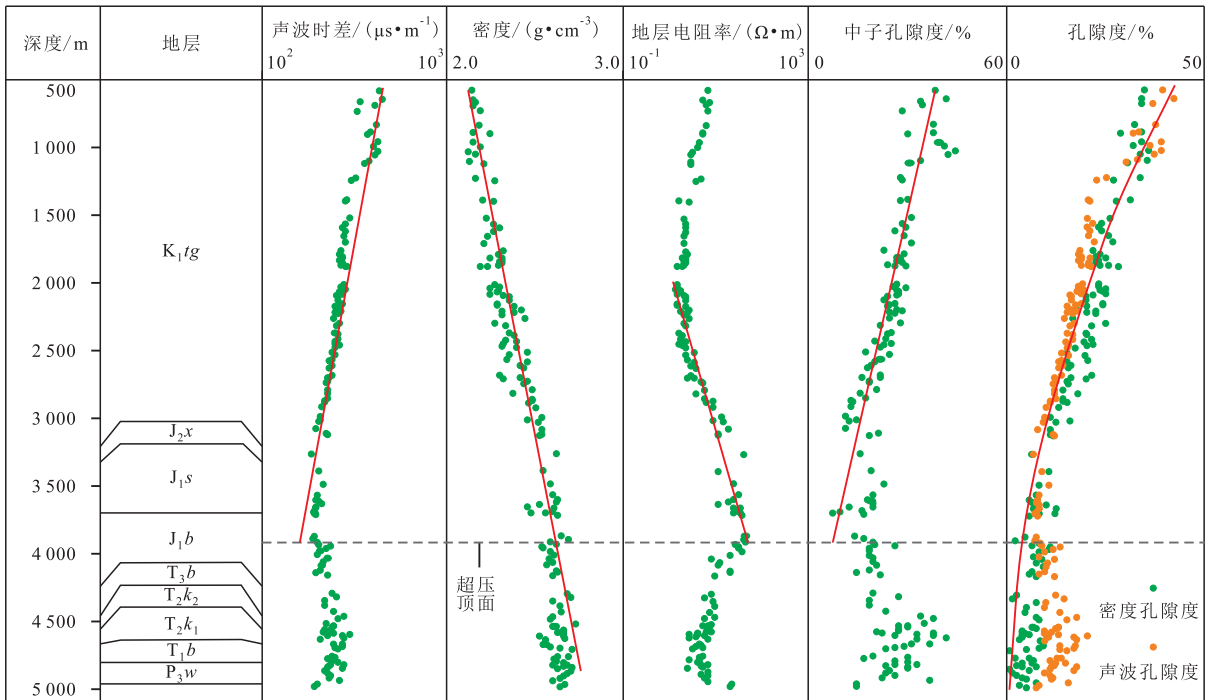


图4 Sx18井测井曲线组合分析
Fig.4 Analysis of well logging curve combination of Well Sx18

由于盆1井西凹陷及周缘位于准噶尔盆地腹部地区,所受到的构造挤压作用并不强烈,因此构造挤压作用对研究区超压形成的影响较小^[18,20]。

3.2 交会图版法

利用声波速度-密度交会图版和垂向有效应力-密度交会图版,结合测井曲线组合分析法对研究区超压进行综合分析。HOESNI研究表明,声波速度-密度交会图版可区分欠压实、构造挤压、成岩作用

及流体膨胀等超压成因^[47]。声波速度与密度在正常压实关系范围之内增加,为欠压实作用;声波速度与密度同时增加,且超过正常压实段为构造挤压,声波速度不变或减小、密度增加或不变为流体膨胀超压或超压传递。

据 Terzaghi 有效应力原理^[48],垂向有效应力是指岩层骨架承受其正上方垂直方向压实作用产生的应力,垂向有效应力可表示为:

$$\sigma = S - p \quad (1)$$

上覆载荷可以通过密度对深度积分得到:

$$S = 10^{-3} \int_0^h \rho g dh \quad (2)$$

针对浅层密度资料不全的问题,可以拟合深度与密度的关系式,通过积分得到上覆载荷与深度的函数^[49]。通常密度随深度呈幂函数变化,据此拟合出Dt1井和Sx18井的密度-深度关系式为:

$$\rho_{Dt1} = 1.292 1h_{Dt1}^{0.079} \quad (3)$$

$$\rho_{Sx18} = 0.977 7h_{Sx18}^{0.1144} \quad (4)$$

通过(1)和(2)式以及2口井拟合得出的密度-深度关系式,可以计算出垂向有效应力,并结合声波速度、密度测井资料,编制声波速度-密度交会图版和密度-垂向有效应力交会图版,进而判断超压形成机制^[40]。通过分析研究区2口典型井Dt1和Sx18的声波速度-密度交会图版(图5a,图6a)可以看出,泥岩段超压表现为声波速度减小,密度变化不明显,表明超压是由生烃膨胀等卸荷增压作用产生的。结合2口井密度-垂向有效应力交会图版(图5b,图6b)发现,其风城组和下乌尔禾组储层砂岩的泥岩超压段邻近泥岩投影点,且偏离正常压实趋势

线,表明存在超压传递增压。结合上述地层信息和地质条件分析认为,超压传递发生的地质条件为早二叠世末隆升期风城组顶部发育的不整合面和中侏罗世由于燕山运动产生的断裂-微裂缝,超压可沿不整合面和断裂-微裂缝分别发生侧向传递和垂向传递。

受构造作用的影响,研究区发育多条断裂,倾角为80°~90°,且受石炭系古凸起的控制,二叠系内部发育局部不整合面^[28,50];其中,早二叠世隆升期风城组顶部出现不整合面^[27],这些二叠系顶面的不整合面在全盆地分布,侧向传导能力强。中侏罗世,由于燕山运动地层在构造抬升作用影响下形成微裂缝;晚侏罗世至早白垩世,风城组烃源岩开始排烃,通过构造作用垂向运移至构造高部位聚集成藏;晚白垩世至第三纪早期,下乌尔禾组烃源岩开始排烃,生成的高成熟油气通过不整合面和断裂带发生运移;至第三纪以后,由于构造活动比较强烈,下乌尔禾组烃源岩产生的高成熟气开始排出。深层、超深层烃源岩由于生烃作用产生的超压会沿着断裂和不整合面从凹陷到隆起、从高势能区向低势能区发生垂向和侧向运移^[51-52](图7),进而在二叠系

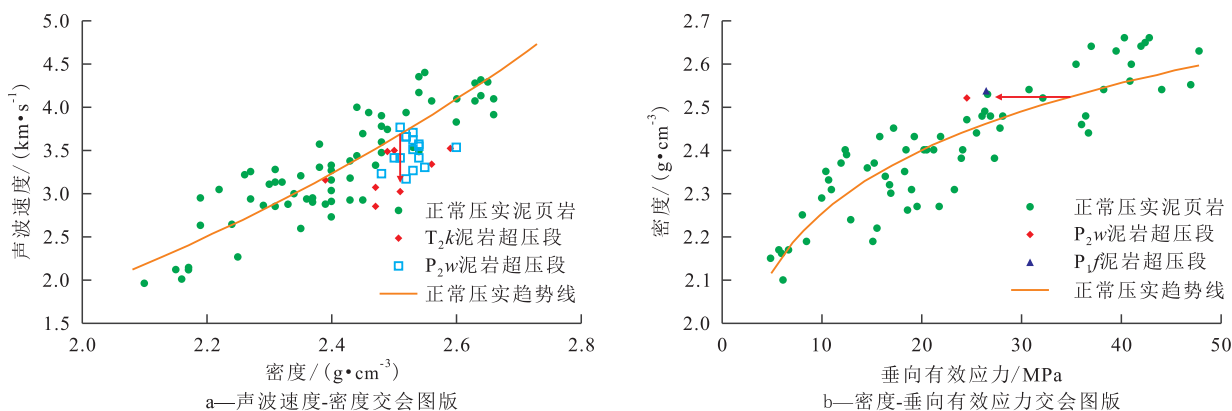


图5 Dt1井交会图版法分析结果
Fig.5 Analysis result of cross plots of Well Dt1

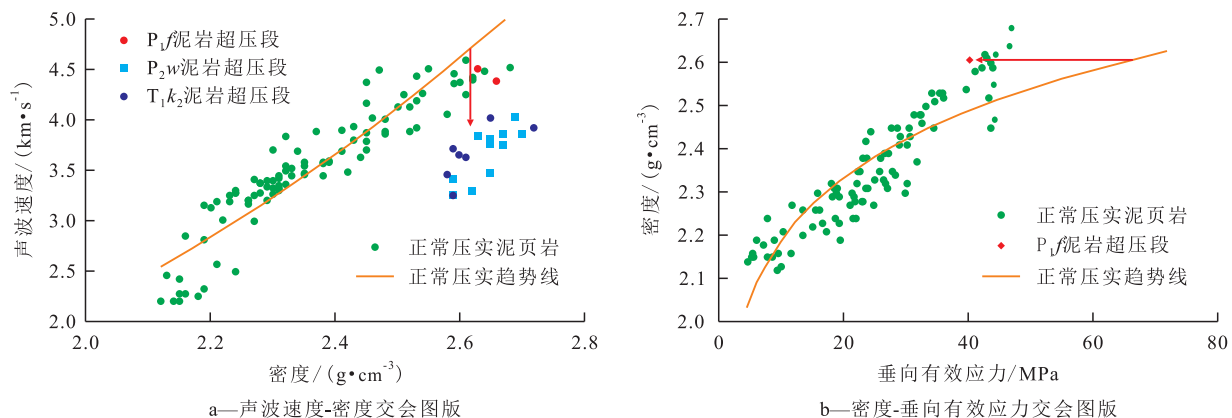


图6 Sx18井交会图版法分析结果
Fig.6 Analysis result of cross plots of Well Sx18

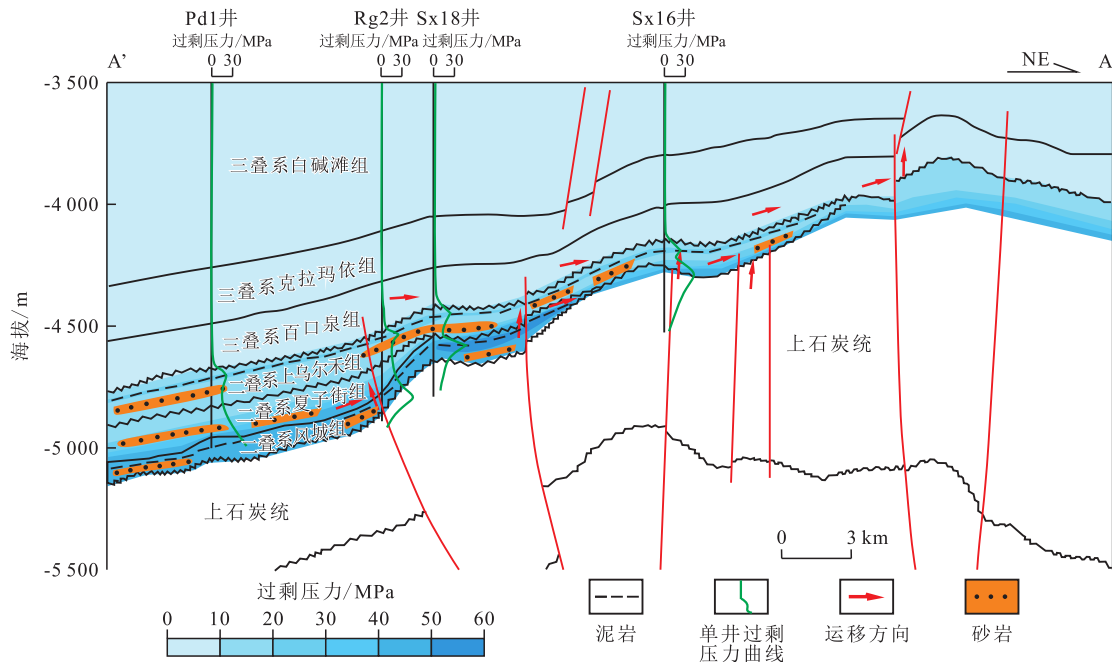


图7 过Pd1井—Sx16井SW—NE向过剩压力剖面分布(剖面位置见图1)

Fig.7 Distribution of southwest-northeast overpressure profile of Wells Pd1-Sx16(See Fig.1)

储层中产生超压^[6]。

超压在传递过程中,密度变化不明显,垂向有效应力明显降低,使超压数据点偏离加载曲线,该点与加载曲线的距离为超压传递增量。通过密度-垂向有效应力交会图版(图5b,图6b)计算出Dt1井风城组和下乌尔禾组储层超压传递增量分别为11.07和9.84 MPa,其占总过剩压力的比例分别为23.35%和21.86%,储层欠压实作用占总过剩压力的比例分别为76.65%和78.14%;Sx18井风城组超压传递增量为25.57 MPa,其储层超压主要来源于深部地层超压沿断裂带向上部地层传递,超压传递贡献率为100%。

综上所述,准噶尔盆地盆1井西凹陷及周缘不同构造部位二叠系岩性地层超压成因不同。达巴松凸起烃源岩层超压成因为欠压实和生烃作用,欠压实和超压传递作用为其储层超压的主要成因,泥岩盖层超压由欠压实作用产生。石西凸起烃源岩层超压由生烃作用产生,储层超压来源于超压传递作用,欠压实作用为泥岩盖层超压成因。

4 超压演化特征

在地质过程分析中,埋藏史、热史、烃源岩生烃过程等可以为超压的产生和保存提供信息^[53-54],通过盆地模拟软件对这些过程进行数值模拟,可以定量评价各成因超压^[22]。采用PressMod软件对地层

压力进行数值模拟,明确研究区目的层超压的形成和演化特征以及定量化各超压机制对现今超压的贡献率^[55],生烃增压评价采用GUO等建立的生烃增压模型^[56-57]。

4.1 数值模拟参数设置

开展流体压力模拟之前,首先对研究区岩性进行精细划分,恢复地层剥蚀量,借鉴李建忠等的研究成果^[37],各井剥蚀量取100~1000 m不等。研究区古热流参考饶松等的研究成果^[58-59],取值为35~52 mW/m²。古地表温度采用PetroMod软件中Auto Swit自动生成温度。Dt1井发育二叠系风城组黑灰色泥岩、白云质泥岩和下乌尔禾组暗色泥岩2套烃源岩,烃源岩的地球化学特征采用钱海涛等的研究成果^[31]。其中,风城组烃源岩干酪根类型为II型,氢指数为3.33~1872.37 mg/g,取值为450 mg/g,总有机碳含量为0.03%~4.43%,取值为1.5%;下乌尔禾组烃源岩干酪根类型为III型,氢指数为1.2~950 mg/g,取值为160 mg/g,总有机碳含量为0.18%~14.03%,取值为1.0%。通过对Dt1井地层、岩性进行划分,以实测温度作为约束条件,对Dt1井埋藏史、热史进行恢复,结果(图8)显示,Dt1井从早三叠世开始进入生油阶段,至侏罗世末期进入生气阶段,现今已达到高成熟-过成熟阶段,主要生成轻质油和天然气。

4.2 烃源岩层超压演化特征

通过模拟Dt1井烃源岩层的过剩压力发现,风

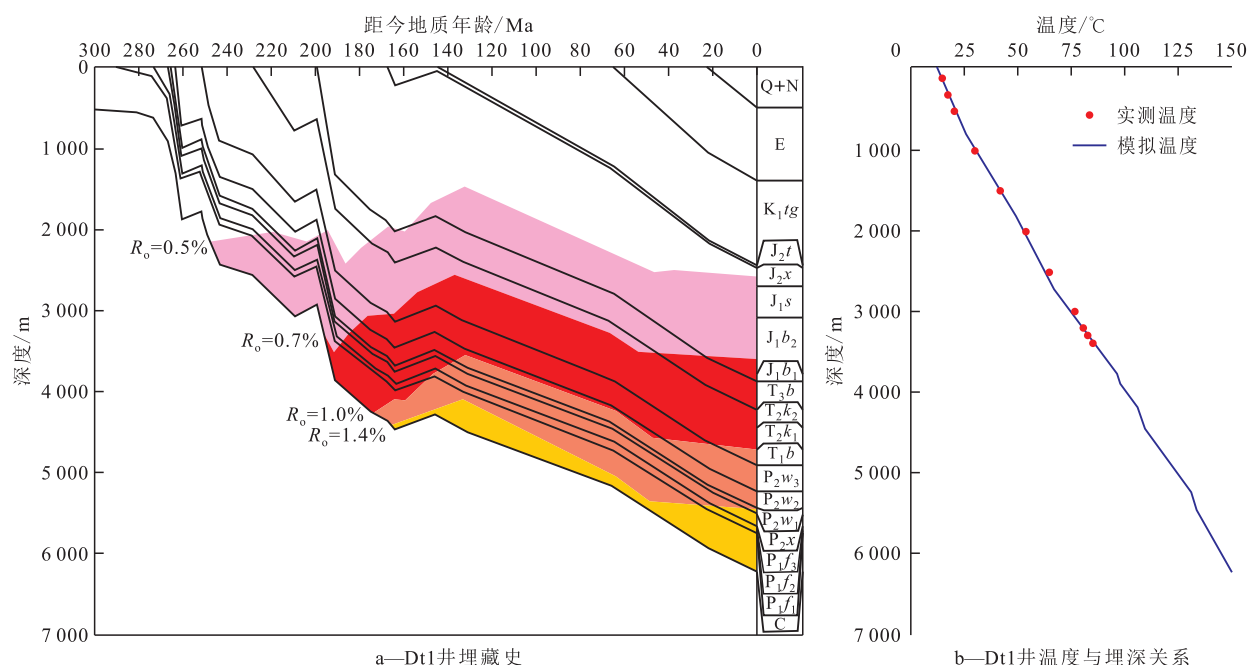


图8 达巴松凸起Dt1井埋藏史和热史恢复

Fig.8 Burial history and thermal history restoration of Well Dt1 in Dabasong Salient

城组烃源岩从距今277 Ma开始生烃,总过剩压力至现今达到最大,为53.89 MPa;现今生烃过剩压力为49.14 MPa,占总过剩压力的91.18%;欠压实增压至现今达到最大,为4.75 MPa;现今地层过剩压力为53.89 MPa,对应地层压力系数为1.94,为强超压(图9a)。其中,从距今200 Ma至170 Ma,生烃增压量突然增大,通过埋藏史和热史分析结果(图8)可以看出,在距今200 Ma开始进入大量生油阶段,导致烃源岩层过剩压力快速增大。下乌尔禾组生烃增压在现今达到最大,为12.48 MPa,现今生烃过剩压力为11.86 MPa,占总过剩压力的84.49%;从距今23 Ma至今,欠压实作用增压从264 Ma开始,缓慢增加,至现今达到最大,为2.18 MPa;下乌尔禾组现今过剩压力为14.04 MPa,对应的地层压力系数为1.25,属于弱超压(图9b)。

Sx18井风城组烃源岩层过剩压力演化大致可以分为3个阶段:第1阶段为距今272 Ma至140 Ma,在距今140 Ma时,过剩压力缓慢增至3.97 MPa,对应的地层压力系数为1.15,为近常压;第2阶段从距今140 Ma至65 Ma,该阶段超压快速增加,至该阶段末期其过剩压力增至41.55 MPa,对应的地层压力系数为1.73,达到强超压;第3阶段从距今65 Ma至今,过剩压力平缓增加,现今为41.76 MPa,对应的地层压力系数为1.83,现今为强超压。生烃增压作用产生的超压在现今达到最大,为39.43 MPa,占总过剩压力的94.41%(图10)。

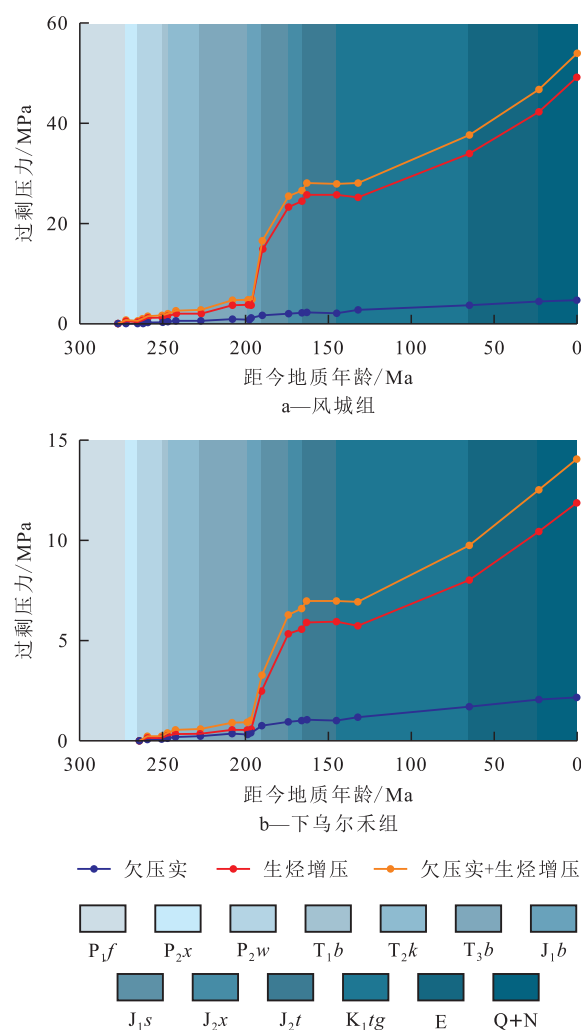


图9 Dt1井烃源岩层过剩压力演化史

Fig.9 Evolution history of overpressure in source rock layers of Well Dt1

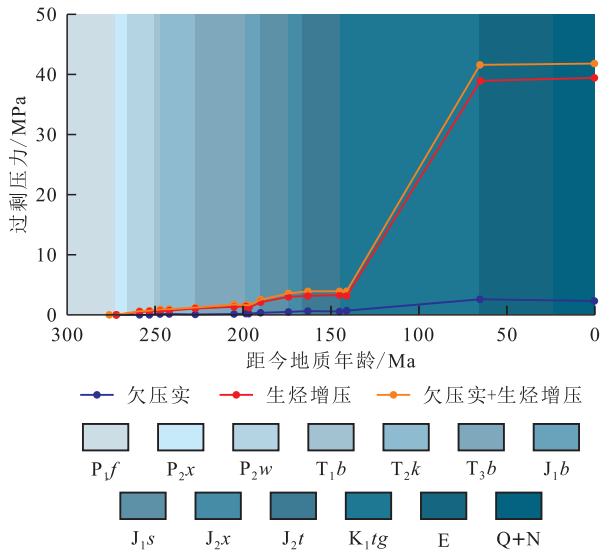


图10 Sx18井风城组烃源岩层过剩压力演化史
Fig.10 Evolution history of overpressure in source rock layers of Well Sx18 in Fengcheng Formation

准噶尔盆地盆1井西凹陷深部是油气聚集的有利区,也是油气的高产区,研究区探井并未钻达凹陷更深处,为了模拟研究区不同构造部位烃源岩过剩压力演化,根据邻井的钻井、测井及地质资料,创建人工井Rg1和Rg2(井位见图1),分别位于盆1井西凹陷和石西凸起(图7)。利用PressMod软件对2口人工井地层发育超压进行数值模拟,并分别用其邻井Dt1和Sx18井的实测温度数据作为约束条件。模拟结果(图11,图12)显示,Rg2井烃源岩层生烃增压量为49.16 MPa,总过剩压力为54.71 MPa,生烃贡献率为89.86%,烃源岩层过剩压力演化与Sx18井的演化过程基本相同;Rg1井风城组和下乌尔禾组烃源岩层生烃增压量分别为44.75和11.77 MPa,总过剩压力分别为68.80和17.49 MPa,生烃贡献率为65%~67.3%,过剩压力演化较Dt1井稍微有所变化。其差异可能是由于越靠近盆地深处,超压顶界深度也随之增大,且Dt1井P_{3w}沿背斜上倾方向尖灭导致地层缺失,其沉积速度过快,过剩压力快速增加,而Rg1井地层沉积比较完整,沉积速度较缓,过剩压力变化较慢。

4.3 盖层超压演化特征

盖层泥岩超压成因主要是由于欠压实作用,随着深度增大,其过剩压力也逐渐变大。在距今200 Ma时,由于快速沉积,导致过剩压力快速增大;在距今166 Ma时,地层缓慢抬升,其过剩压力又逐渐变小,后又随深度增大,从距今145 Ma逐渐增至现今的11.3 MPa(图13a)。Sx18井盖层变化在距今65 Ma以前基本与Dt1井类似,从距今65 Ma至今,

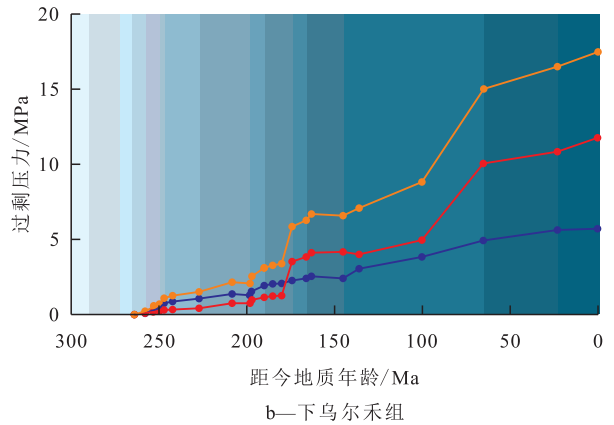
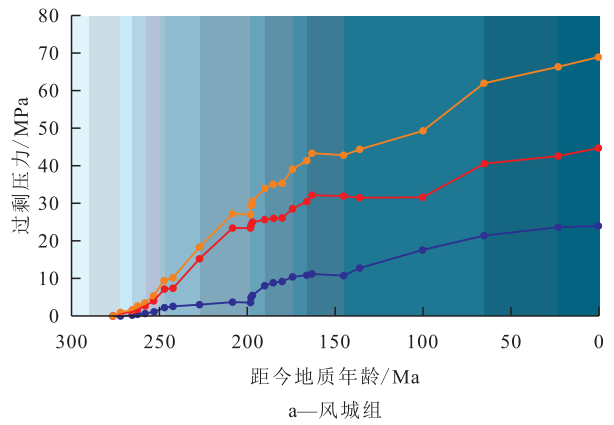


图11 Rg1井烃源岩层过剩压力演化史
Fig.11 Evolution history of overpressure in source rock layers of Well Rg1

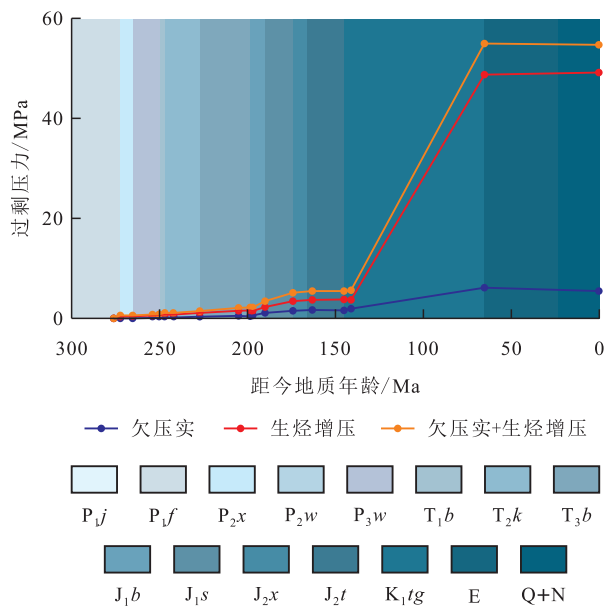


图12 Rg2井烃源岩层过剩压力演化史
Fig.12 Evolution history of overpressure in source rock layers of Well Rg2

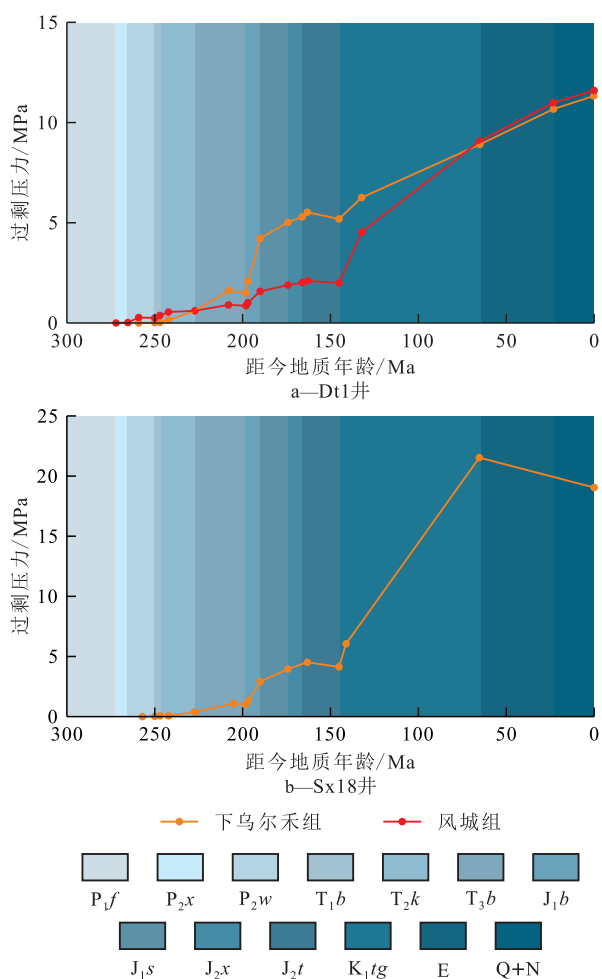


图13 Dt1井和Sx18井盖层过剩压力演化史
Fig.13 Evolution history of overpressure in caprock of Wells Dt1 and Sx18

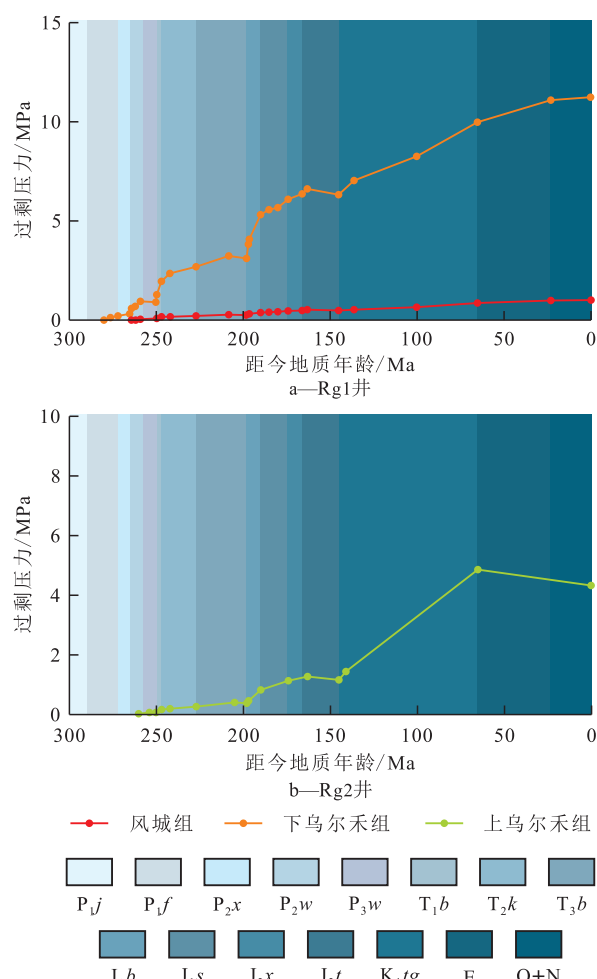


图14 Rg1井和Rg2井盖层过剩压力演化史
Fig.14 Evolution history of overpressure in caprock of wells Rg1 and Rg2

由于其地层一直抬升,过剩压力逐渐变小,现今为19.06 MPa(图13b)。盖层泥岩过剩压力为1.03~19.06 MPa(图13a,图14)。

4.4 储层超压演化特征

超压传递作用也是准噶尔盆地盆1井西凹陷及周缘二叠系储层超压的成因。由图15可以看出,达巴松凸起Dt1井风城组和下乌尔禾组在三叠世末期欠压实作用已经基本形成。中侏罗世燕山运动时期,盆1井西凹陷发生构造抬升作用,形成的断裂与海西运动时期的断裂垂向搭接(图7),风城组烃源岩在晚侏罗世—早白垩世大量生气^[34,51,60],其凹陷区生烃产生的过剩压力达到40.50 MPa,加之欠压实作用产生的过剩压力为21.46 MPa,其总过剩压力达到61.96 MPa,形成了强超压;凸起区生烃产生的过剩压力为33.93~48.71 MPa,加之欠压实作用产生的过剩压力为3.77~6.22 MPa,其总过剩压力达到37.70~54.93 MPa。另外,晚侏罗世至早白垩世为下乌尔禾组烃源岩大量生油气阶段,其凹陷区

该烃源岩生烃产生的过剩压力为10.06 MPa,加之欠压实作用产生的过剩压力为4.95 MPa,总过剩压力达到15.01 MPa;凸起区该烃源岩生烃产生了8.04 MPa的过剩压力,加之欠压实作用产生的过剩压力为1.71 MPa,总过剩压力达到9.75 MPa。同时前人研究证实晚侏罗世至早白垩世也是这2套烃源岩向二叠系储层的主要油气充注时期^[34,51,60],烃源岩层中高过剩压力推动油气通过不整合面和前期断裂分别发生侧向和垂向传递运移,从而导致凸起区风城组和下乌尔禾组储层过剩压力增大(图7);古近纪喜马拉雅运动时期,盆地整体发生由南向北的掀斜运动,生成的高过剩压力致使油气由凹陷向凸起侧向传递运移,导致高部位凸起区风城组和下乌尔禾组储层过剩压力进一步增大,油气聚集藏,超压此时发生小部分传递。综合上述超压传递增压过程分析,认为风城组和下乌尔禾组储层大量超压传递增压主要发生于晚侏罗世至早白垩世,少量发生于古近纪喜马拉雅运动时期,因此,为了定

量化表征其超压传递增压的大小,将晚侏罗世至早白垩世超压传递增压量劈分为现今的90%,为8.86~9.97 MPa,古近纪喜马拉雅运动时期超压传递增压量劈分为现今的10%,为0.98~1.21 MPa(图15)。

准噶尔盆地盆1井西凹陷及周缘凸起区和凹陷区的烃源岩生烃增压均从早二叠世开始,现今达到最大,该增压作用对于凸起区和凹陷区超压形成的贡献率分别为84.49%~94.41%和65%~67.3%。凹陷区和凸起区总过剩压力分别为68.8和41.76~54.71 MPa,其凹陷区总过剩压力较凸起区大。对比不同构造部位风城组和下乌尔禾组总过剩压力发现,其总过剩压力分别为14.04~17.49和41.76~68.8 MPa,风城组总过剩压力普遍大于下乌尔禾组,原因是随着深度的增大,欠压实作用对于深层的影响越来越小,生烃增压作用对于更深层过剩压力影响变大。

超压传递作用对于研究区凸起区二叠系储层

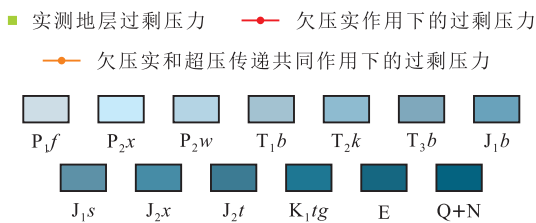
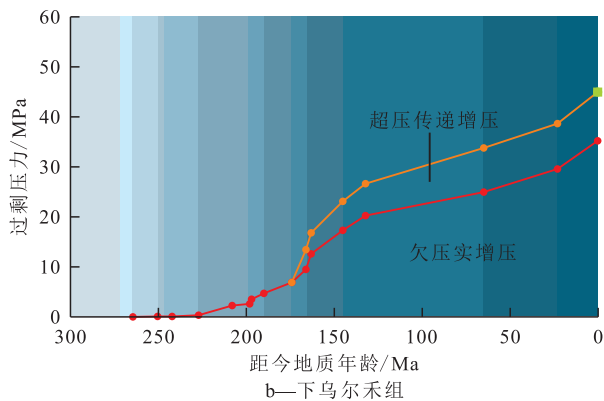
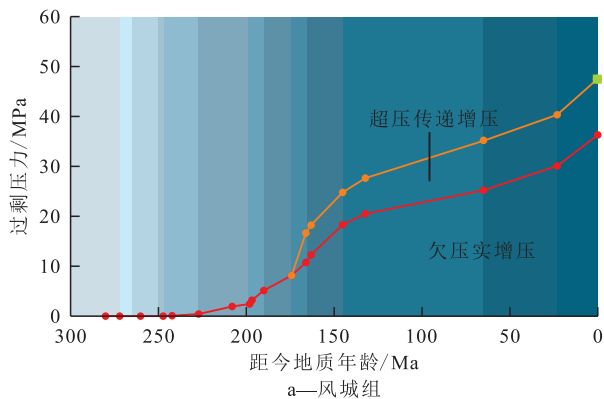


图15 Dt1井不同成因类型储层过剩压力演化特征

Fig.15 Evolution characteristics of overpressure of different genesis types of Well Dt1 reservoirs

超压贡献率存在差异,其中达巴松凸起风城组和下乌尔禾组超压传递贡献率为21.86%~23.35%,石西凸起风城组超压传递贡献率为100%。泥岩盖层过剩压力来源于欠压实作用,过剩压力为1.03~19.06 MPa。

4.5 不同层系超压差异性演化对油气成藏的影响

依据研究区二叠系不同构造部位不同岩性各超压成因机制的判识、演化以及定量评价结果,结合盆1井西凹陷油气关键成藏期分析认为,晚三叠世欠压实作用在风城组和下乌尔禾组储层中基本形成,抑制了储层的压实作用,使储层中的部分原生孔隙得以保存;中侏罗世,燕山运动使断裂开启,促进了储层超压传递作用的产生,抑制了演化后期的压实作用,因此研究区二叠系风城组和下乌尔禾组储层中各类机制超压的产生一定程度上保存了储集空间,改善了储集物性;晚侏罗世,风城组和下乌尔禾组烃源岩与其上覆储层的源储过剩压力差在达巴松凸起为5~38 MPa,石西凸起为2~20 MPa,较大的源储过剩压力差为油气的充注提供了强劲的运移动力条件,评价得到的达巴松凸起风城组和下乌尔禾组储层超压传递增压多为9~11 MPa,也进一步证实了该作用的发生。研究区二叠系泥岩盖层超压在达巴松凸起和石西凸起为11~19 MPa,其可为圈闭中油气提供超压封闭条件,进一步增强盖层的封闭能力。因此,盆1井西凹陷及周缘不同岩性地层超压差异性演化的研究对于油气成藏具有重要意义。

5 结论

准噶尔盆地盆1井西凹陷及周缘超压发育,不同构造部位超压大小存在差异,石西凸起地层超压为17.61~23.58 MPa,属于弱超压;达巴松凸起地层超压为41.20~47.41 MPa,属于强超压。强超压作用下,渗透率增大,深层岩层有形成大规模储层的可能。

研究区深层二叠系超压成因较为复杂,同一地层的不同构造部位和同一位置的不同地层其超压成因存在较大差异。达巴松凸起烃源岩层超压成因为欠压实和生烃作用,储层超压的主要成因为欠压实和超压传递作用,盖层超压由欠压实作用产生;石西凸起烃源岩层超压成因为生烃作用,储层超压成因为超压传递作用,而泥岩盖层超压成因为欠压实作用。源、储、盖层超压成因的研

究对于油气储集、保存和运移具有重要作用。

准噶尔盆地1井西凹陷及周缘不同构造部位的超压贡献率存在差异,研究区凸起区和凹陷区烃源岩层生烃增压自早二叠世开始,至今达到最大,生烃增压贡献率分别为84.49%~94.41%和65%~67.3%。凸起区生烃贡献率大于凹陷区,但凹陷区总过剩压力大于凸起区,并且风城组总过剩压力大于下乌尔禾组。二叠系储层超压传递作用引起超压的贡献率在不同凸起区有较大差异,达巴松凸起为21.86%~23.35%,石西凸起为100%。欠压实作用对研究区泥岩盖层超压的贡献率基本为100%。

符号解释

- g —— 重力加速度, m/s^2 ;
 h —— 深度, m ;
 h_{Dt1} —— Dt1井的深度, m ;
 h_{Sx18} —— Sx18井的深度, m ;
 p —— 流体压力, MPa ;
 S —— 上覆载荷, MPa ;
 ρ —— 地层密度, g/cm^3 ;
 ρ_{Dt1} —— Dt1井的地层密度, g/cm^3 ;
 ρ_{Sx18} —— Sx18井的地层密度, g/cm^3 ;
 σ —— 垂向有效应力, MPa 。

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